

# University of Wyoming **King Air** and **Cloud Radar**. **Airborne millimeter radar research since 1992.**

**Marine stratus off the coast of Oregon (1995, 1999)** ---- *J. Atmos. Sci.*, **55**, 3540-3564, 1998.

**Small cumulus, Florida (1995)** ---- *Atmos. Res.* **52**, 143-165, 1999  
*Q. J. Roy. Meteor. Soc.*, **126**, 415-443, 2000.

**Winter Ns, Ac lent., ... , Wyoming (1992, 1997, 2000)** ---- to appear *J. Atmos. Sci.*, Apr. 2001

**Marine Sc, E. Pacific (DYCOMS, 2001)** ---- coming in July 2001

**URL : [//www-das.uwyo.edu/wcr](http://www-das.uwyo.edu/wcr)**

Pending proposal : “***Improvements to spaceborne cloud profiling, based on aircraft and radar observations of cloud structure.***”

- insights into cloud processes of relevance to CloudSat
- improvements to algorithms for cloud profiling

Pixel size:

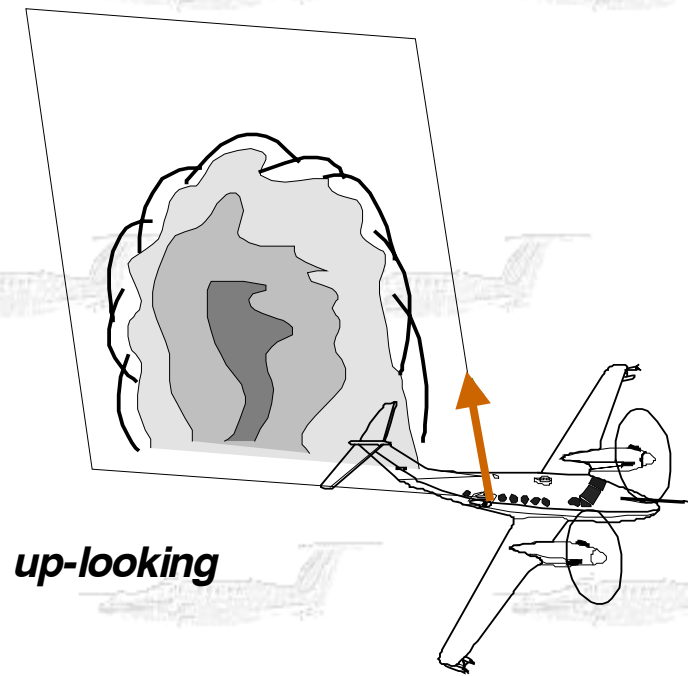
WCR:  $10^2 \times \pi \times 30 \approx 10^4 \text{ m}^3$

CloudSat:  $3 \times 10^9 \text{ m}^3$

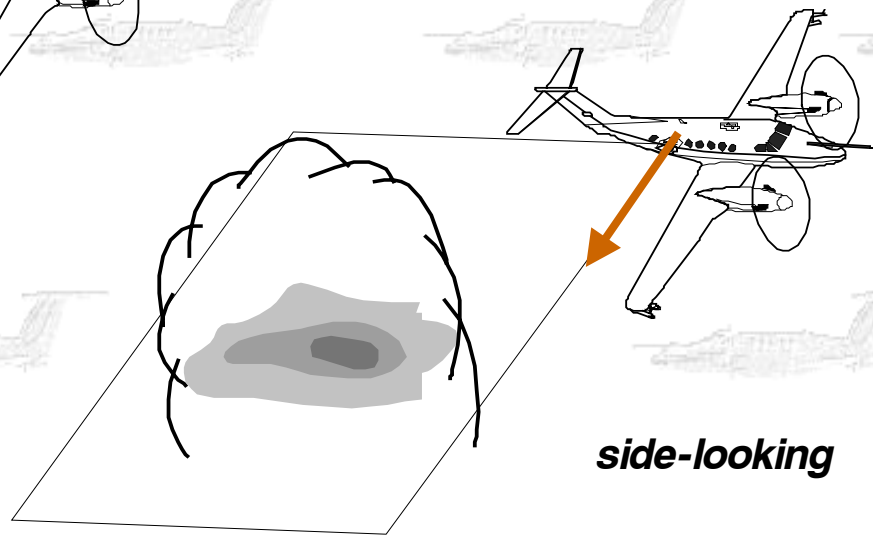
Time to cover 50 km:

WCR: **10 minutes**

CloudSat: **7.5 seconds**



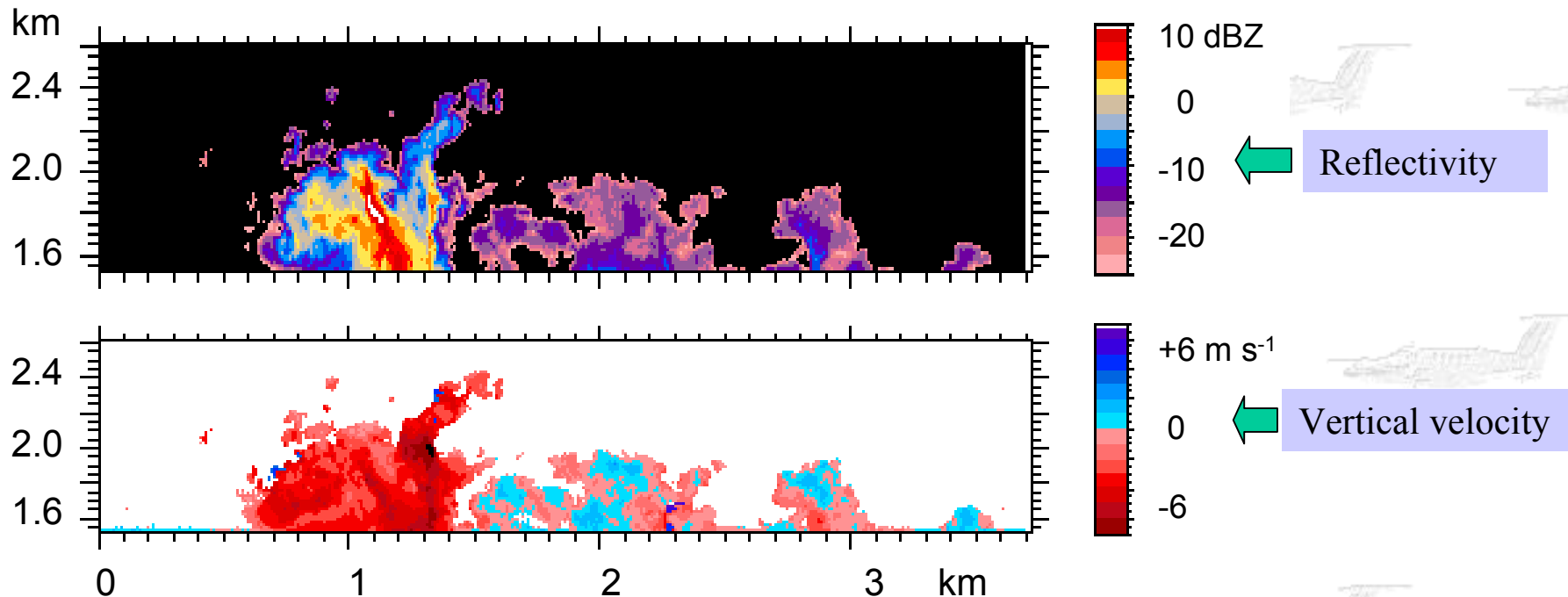
***up-looking***



***side-looking***

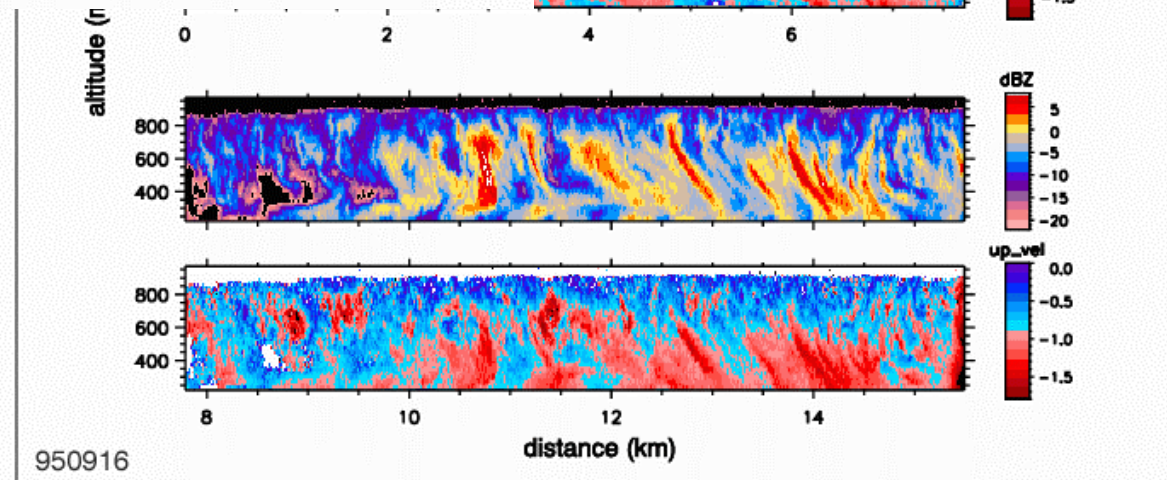
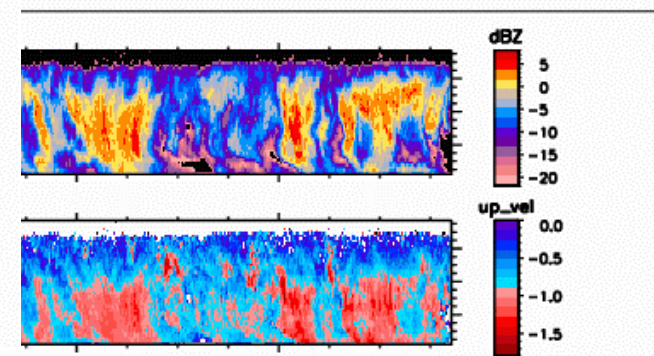
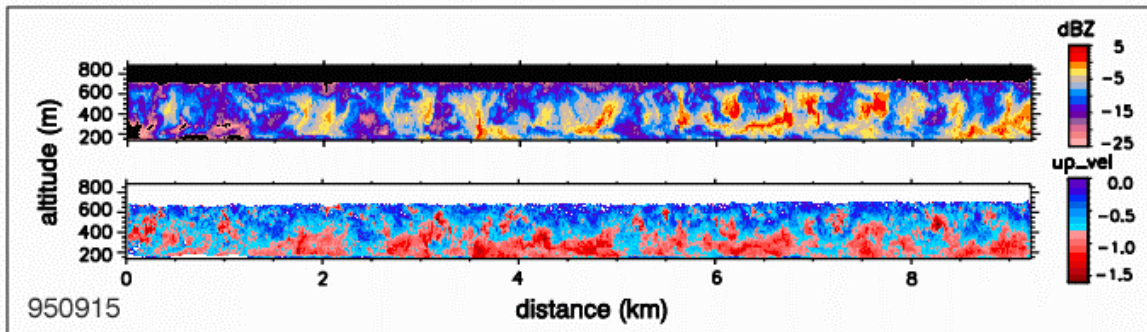
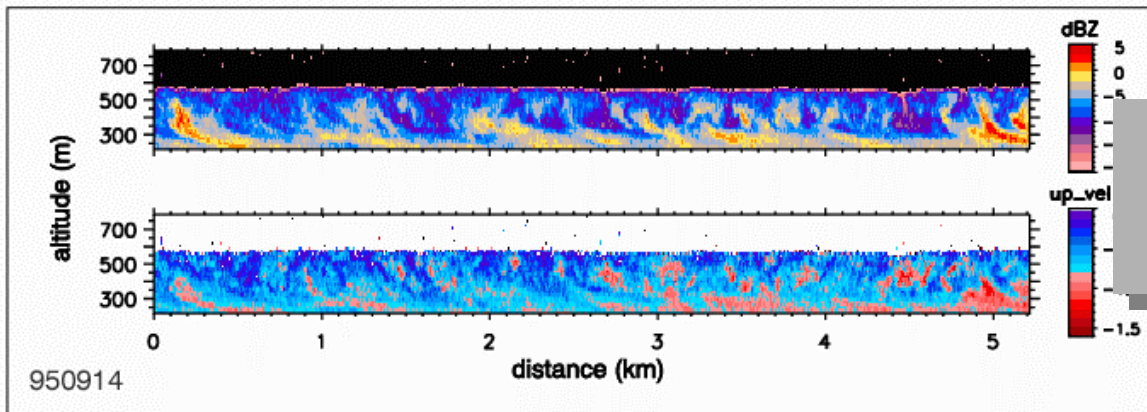
# Cumulus in dissipating stage

Florida, 10 August 1995

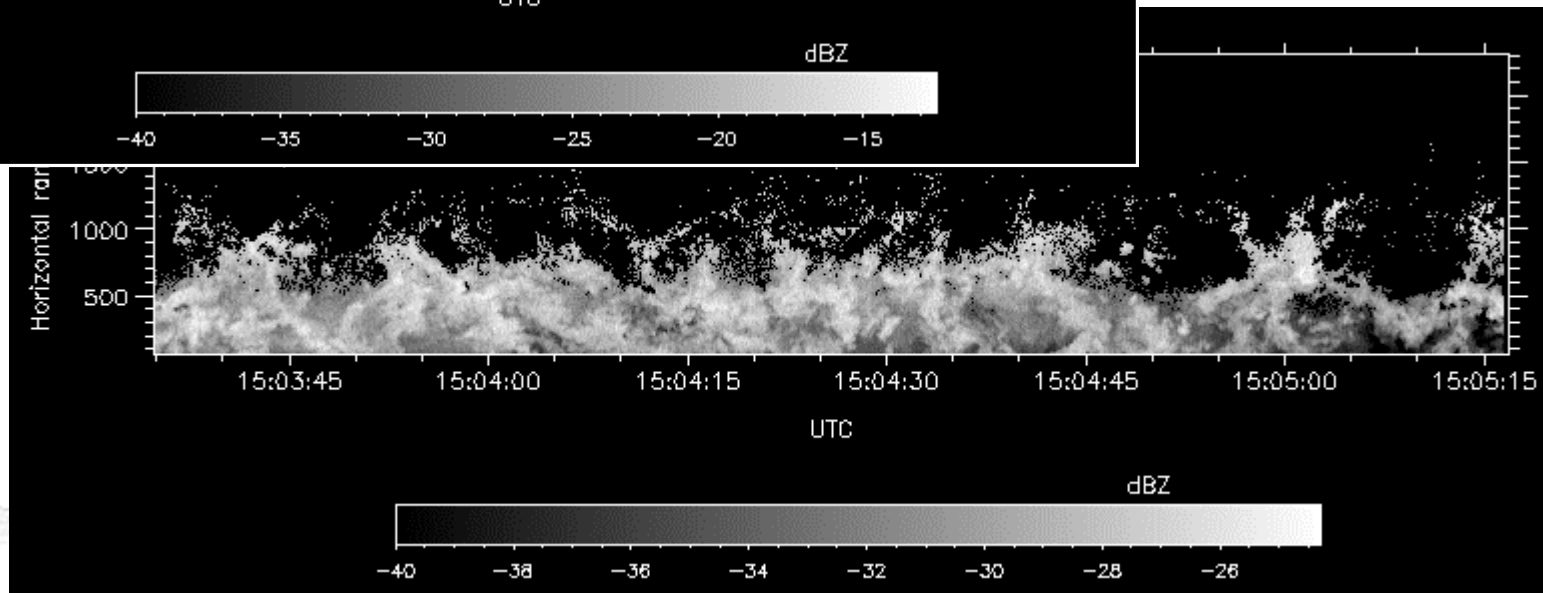
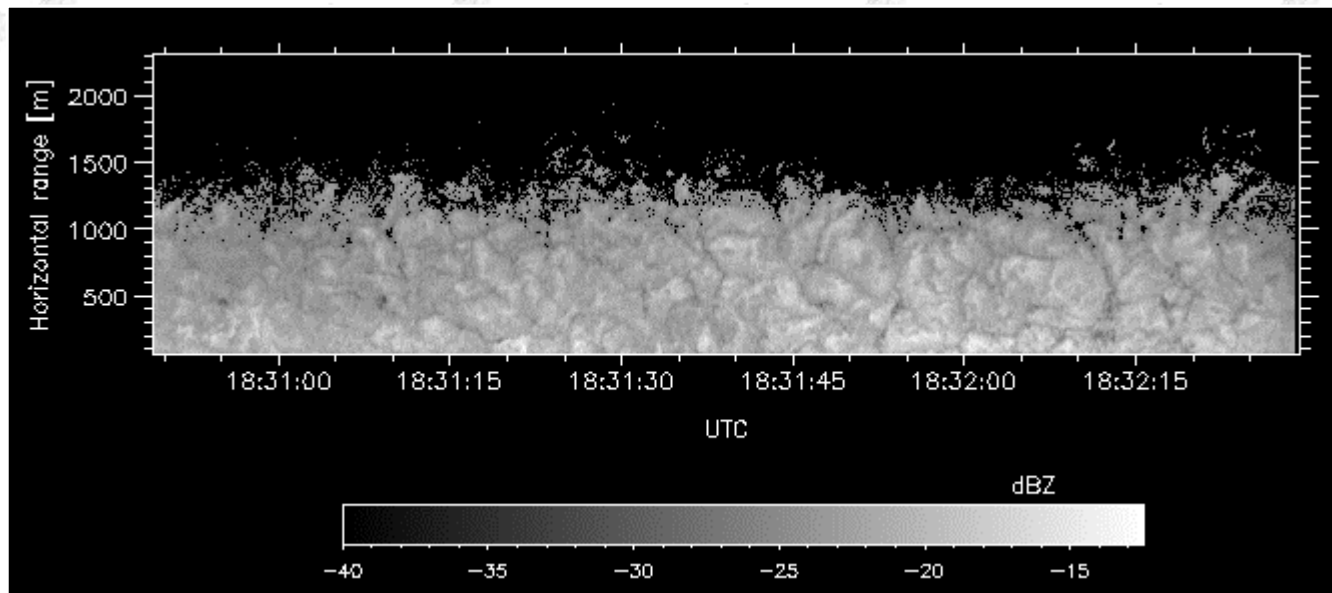


The cloud located at 1 km has reached the dissipating stage, with all of the cloud volume moving downward. At the flight level, the temperature was 18.5°C, the liquid water content 0.7 g m<sup>-3</sup> and very light drizzle with drops 200  $\mu$ m diameter was recorded. The total cloud droplet concentration was about 150 cm<sup>-3</sup>. The reflectivity field indicates the presence of 3 precipitation shafts above 1.8 km altitude. These merge into one main core of precipitation. Smaller clouds at 2 and 2.9 km have some weak updraft regions.

## Coastal stratus Oregon, 1995



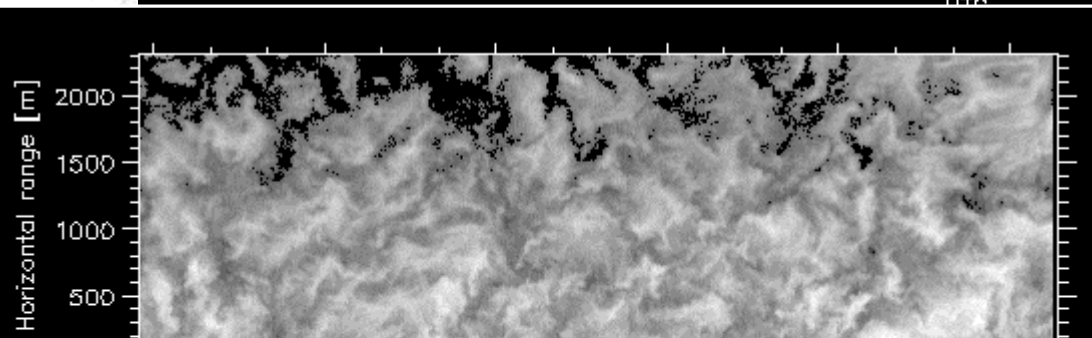
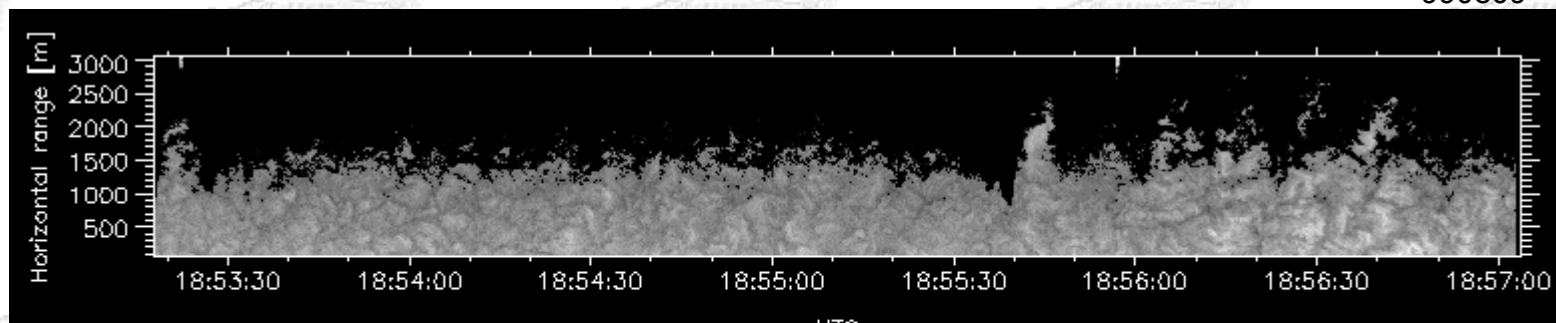
Vertical sections of reflectivity and of vertical velocity. The horizontal and vertical scales of the images are 1:1.



**Horizontal sections of various coastal stratus cases.** Attenuation correction has been applied, but signal to noise ratio limits the range of detection. These images demonstrate large variety in the scales and textures of reflectivity fields.

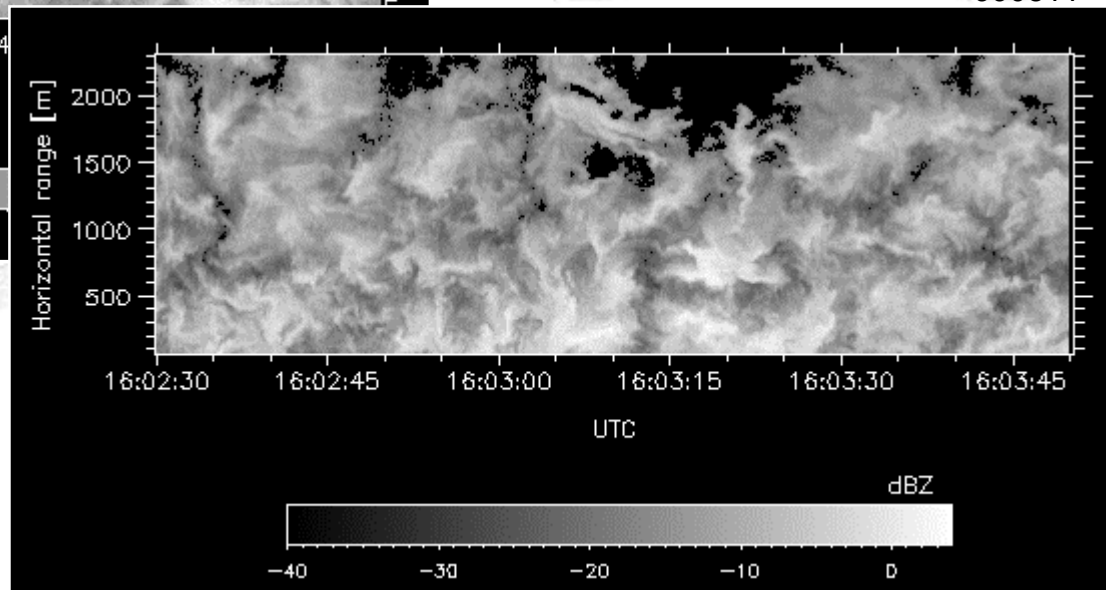


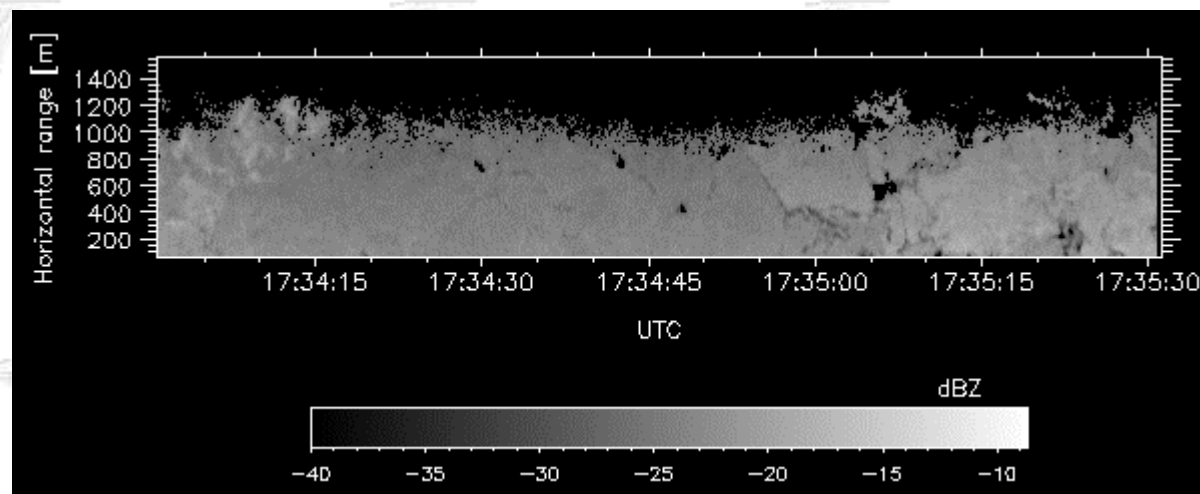
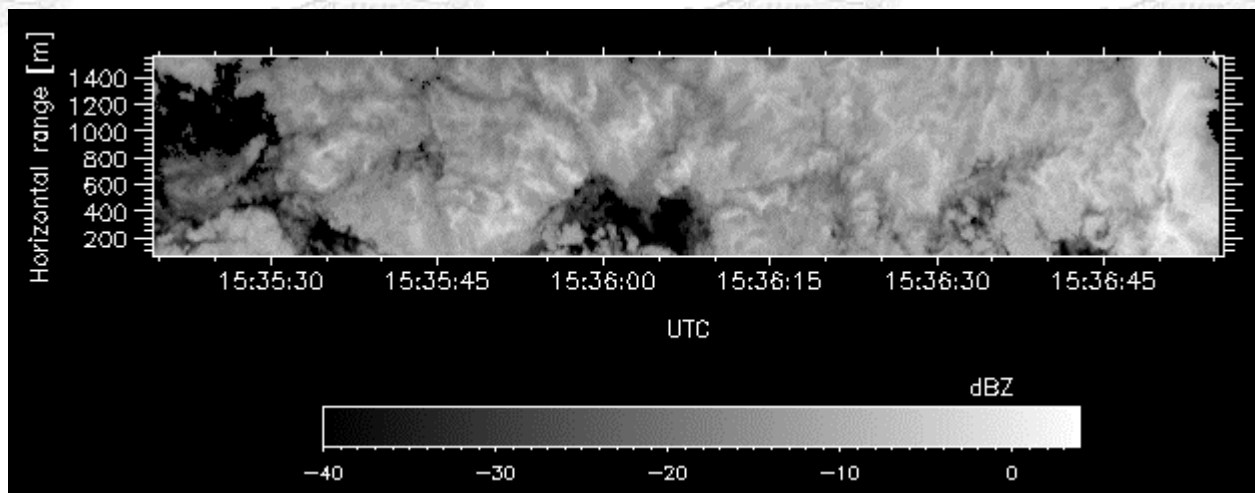
990809



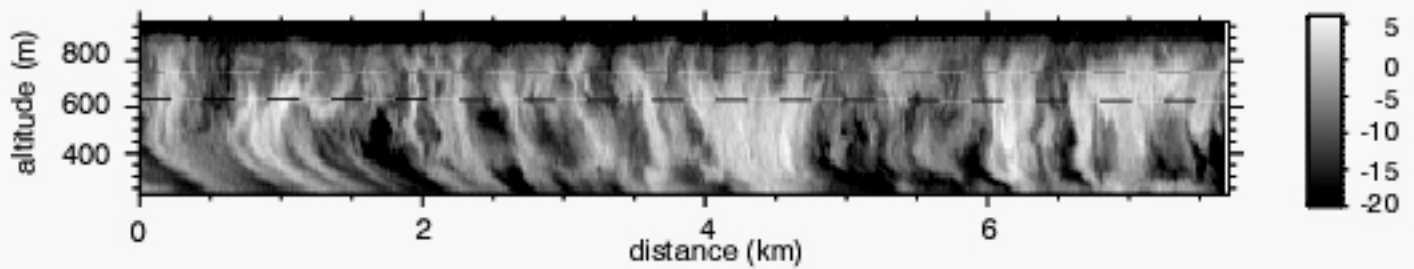
990811

990811

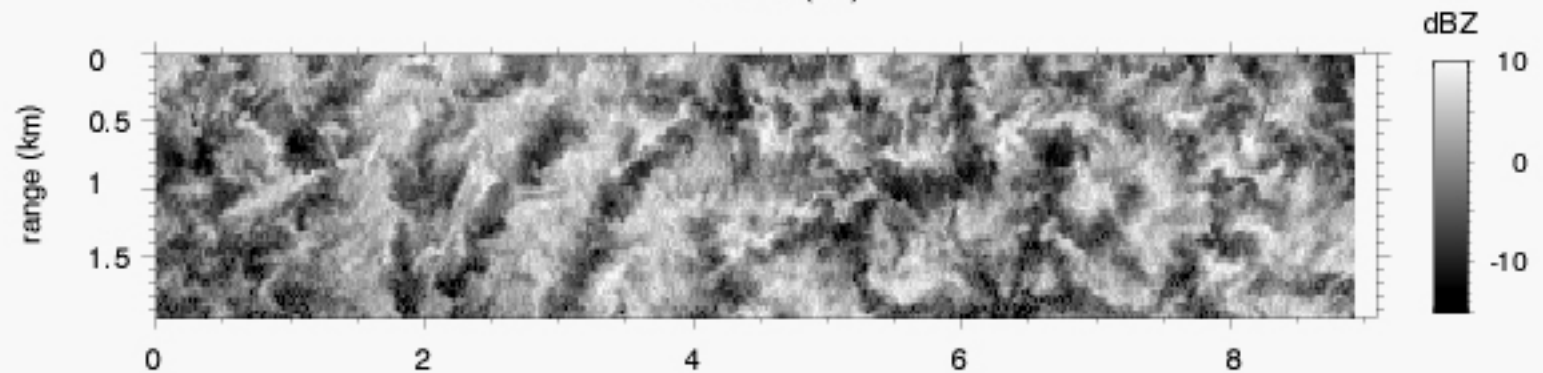
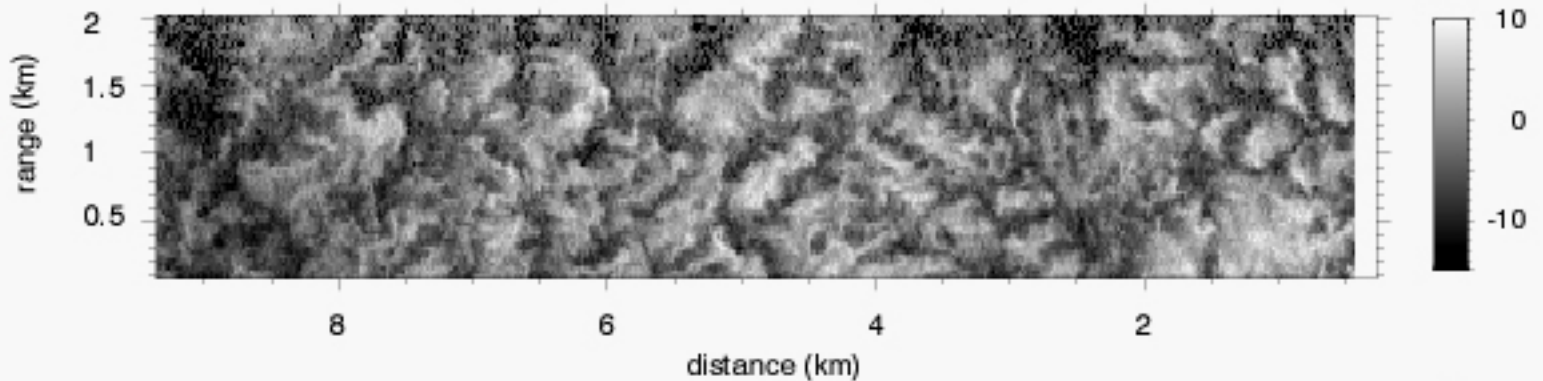




### 950916 - VERTICAL SECTION



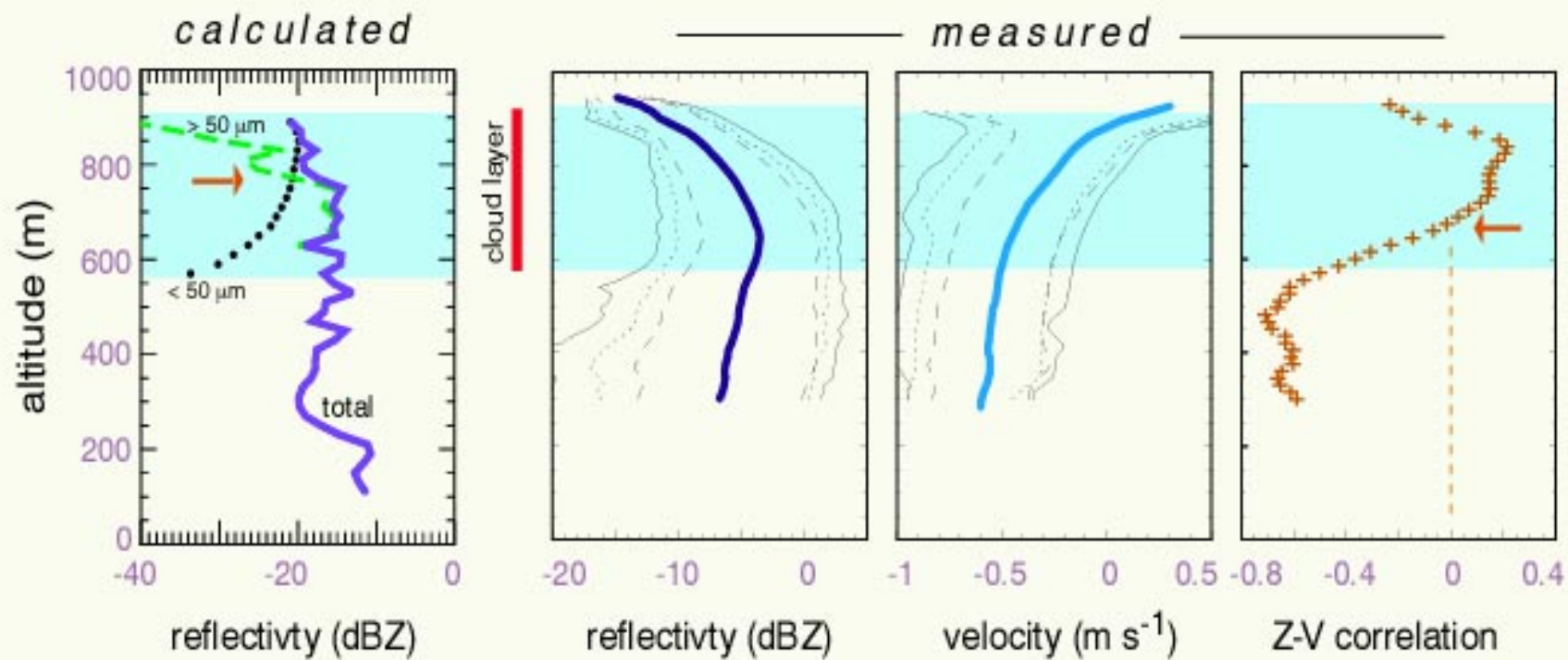
### HORIZONTAL SECTIONS



625 m altitude

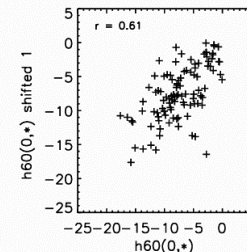
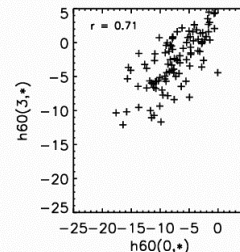
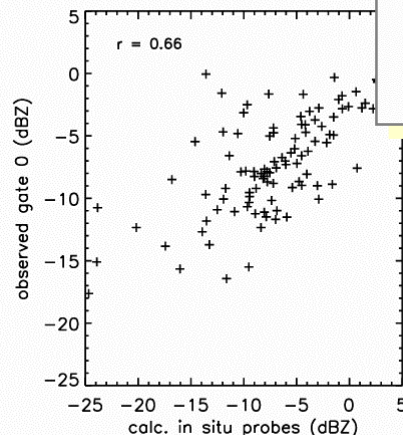
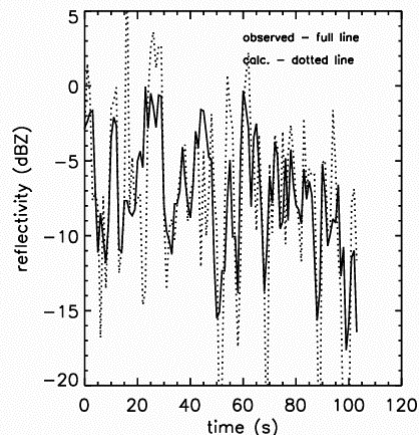
750 m altitude





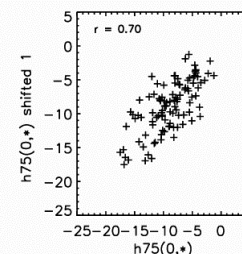
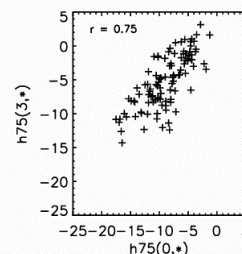
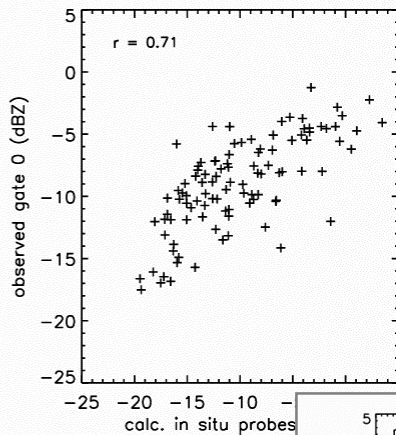
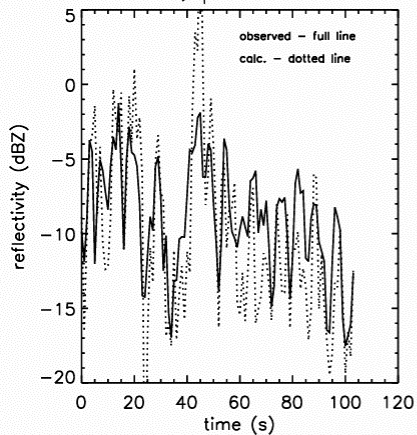
950916

630 m altitude;  $\phi = 0.15$

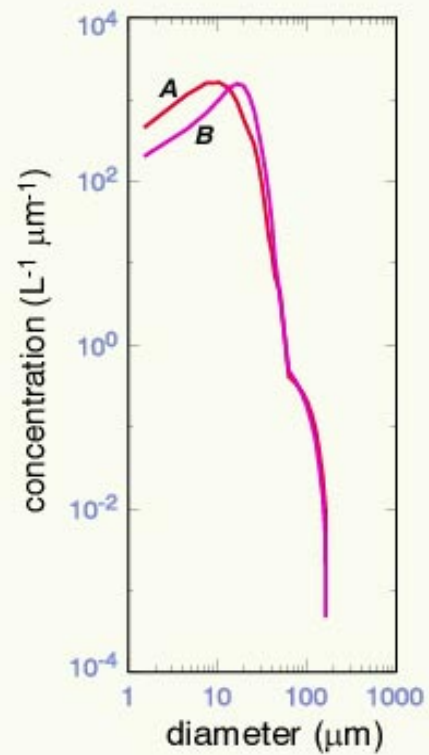
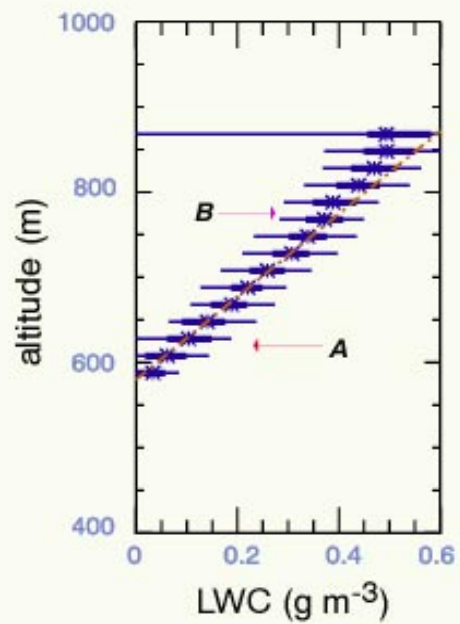


The correlation between  $Z_{calc}$  (derived from in situ data along the flight path) and  $Z_{obs}$  (radar measurement 90 m to the side of the aircraft), is very similar to that between radar ranges separated by a similar distance.

780 m altitude;  $\phi = 0.6$



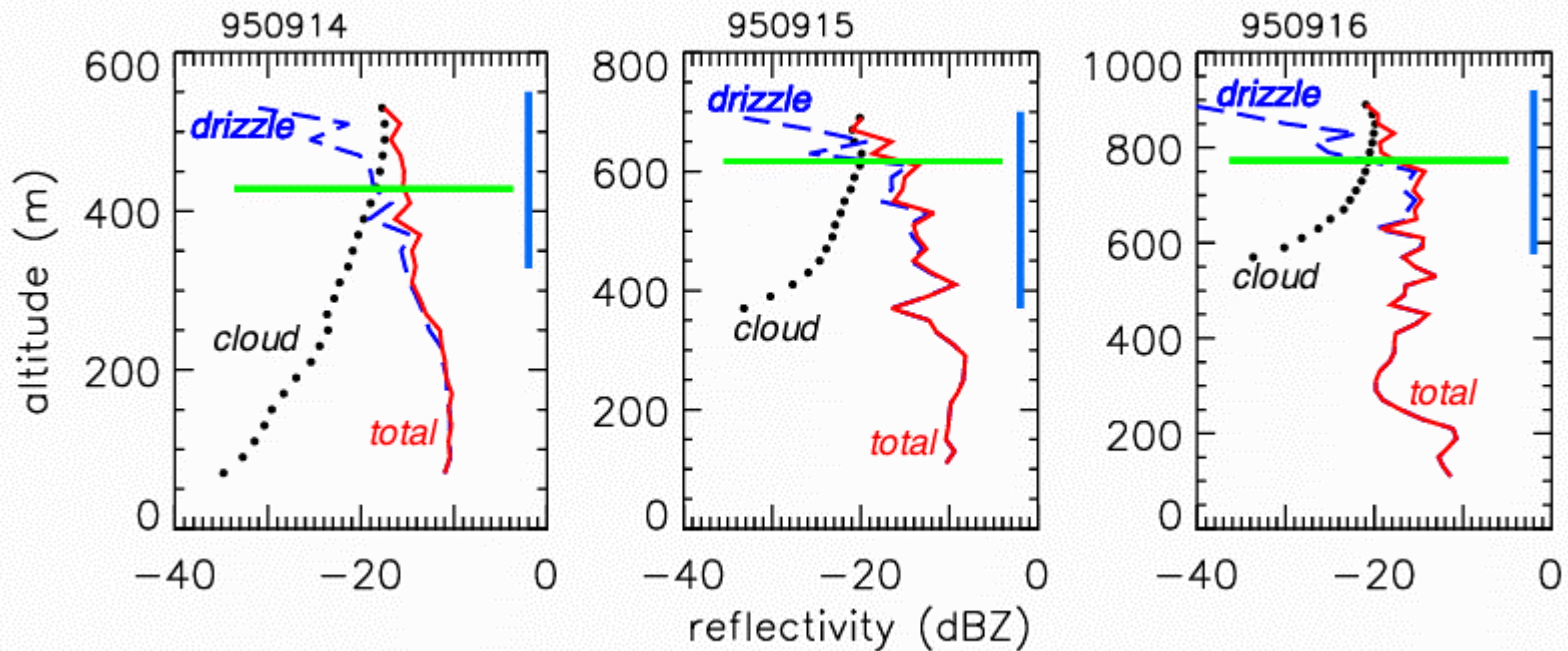
*Interpretation of reflectivity patterns in terms of spectral characteristics is possible.*



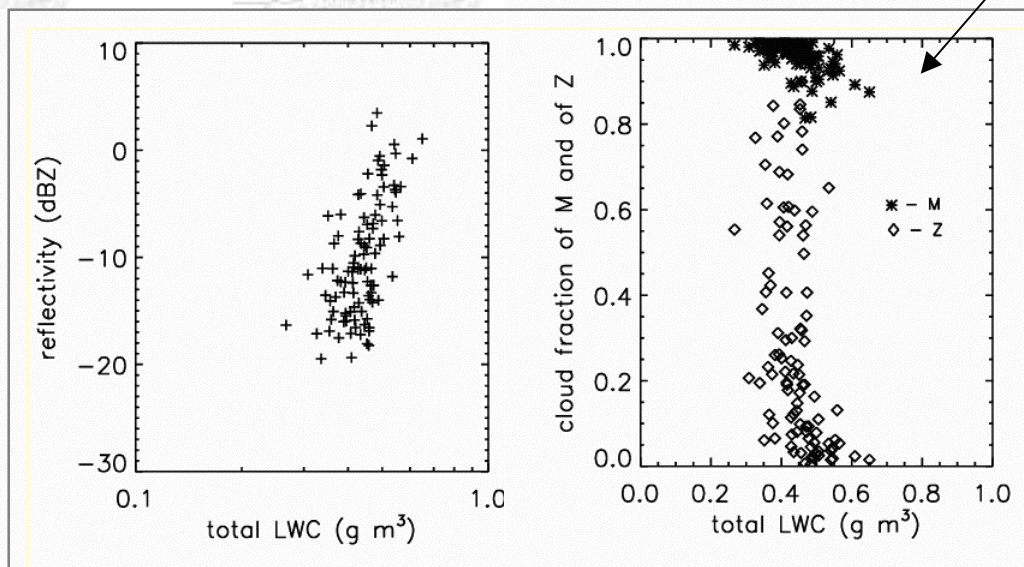
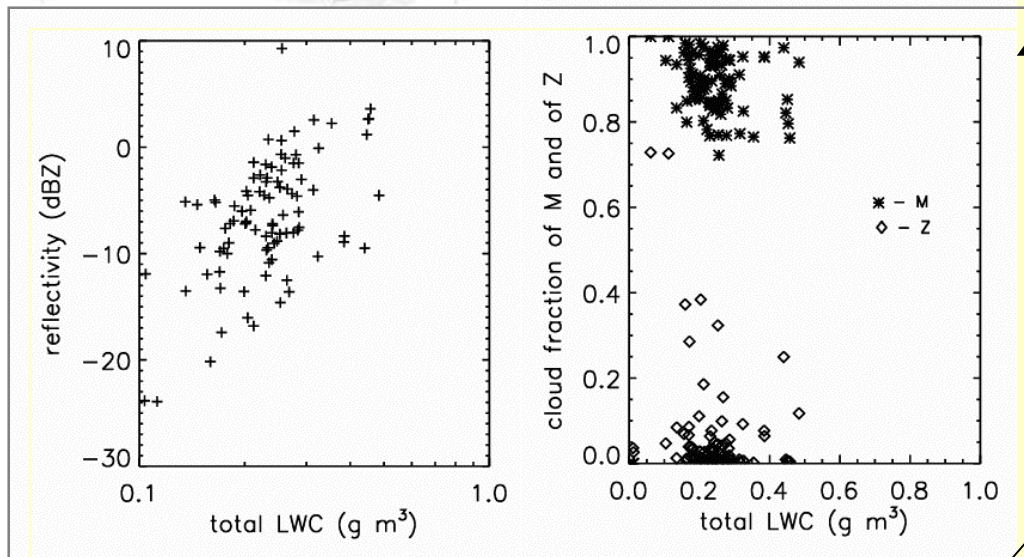
950916

Conclusion from observations of Oregon coastal stratus:

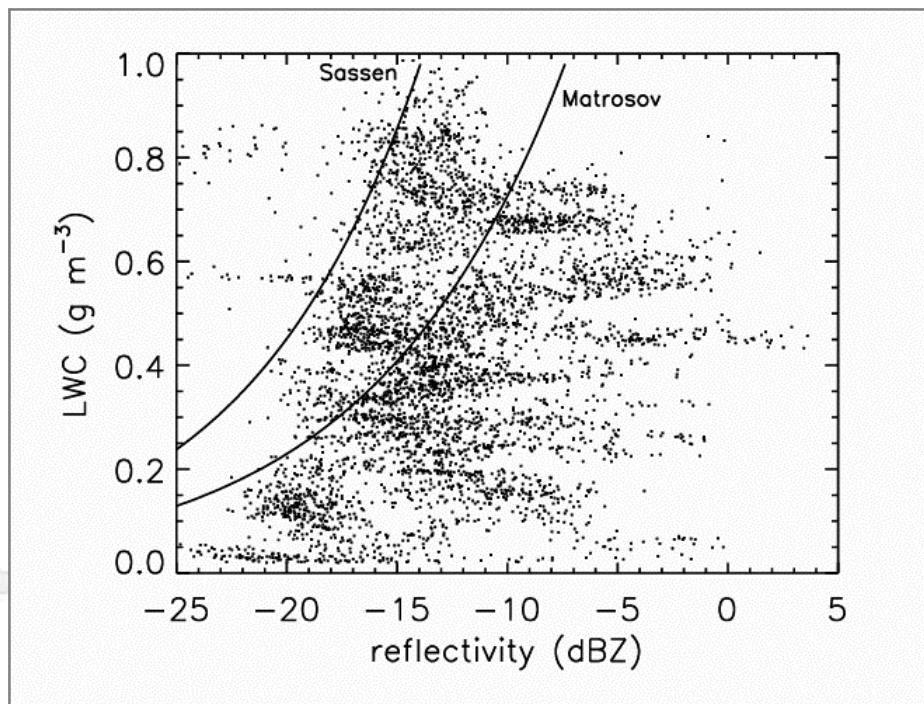
**Reflectivity is dominated by drizzle through the major part of the cloud depth.**







**Mass - reflectivity (Z - M) relationship is complex. Simple rule would be unreliable.**



**The use of M - Z relationships based on cloud droplet spectra, without the inclusion of drizzle, will seriously overestimate LWC from radar data.**

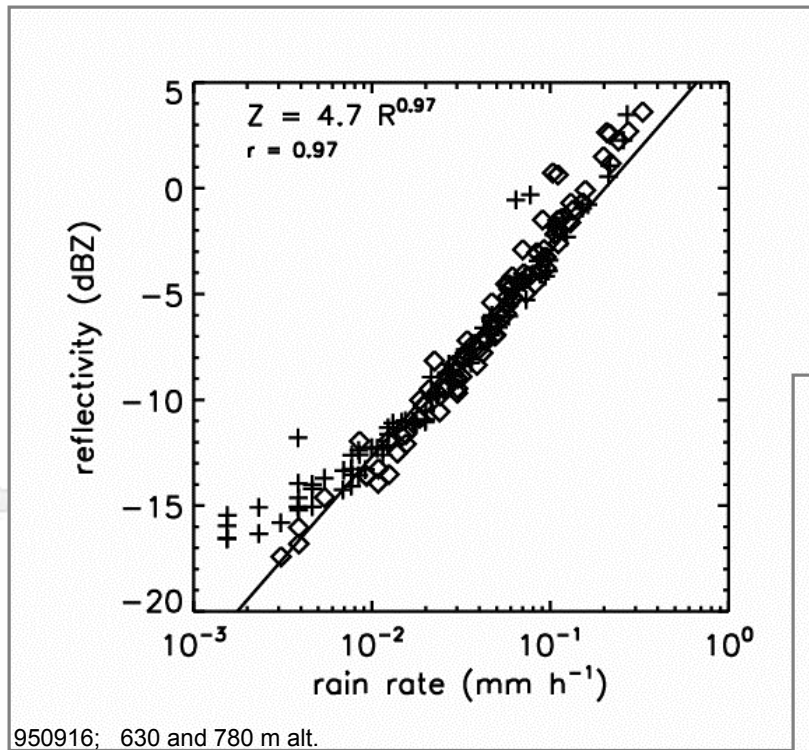
990817 flight

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**Attenuation correction is quite straightforward when mass and reflectivity are decoupled.**



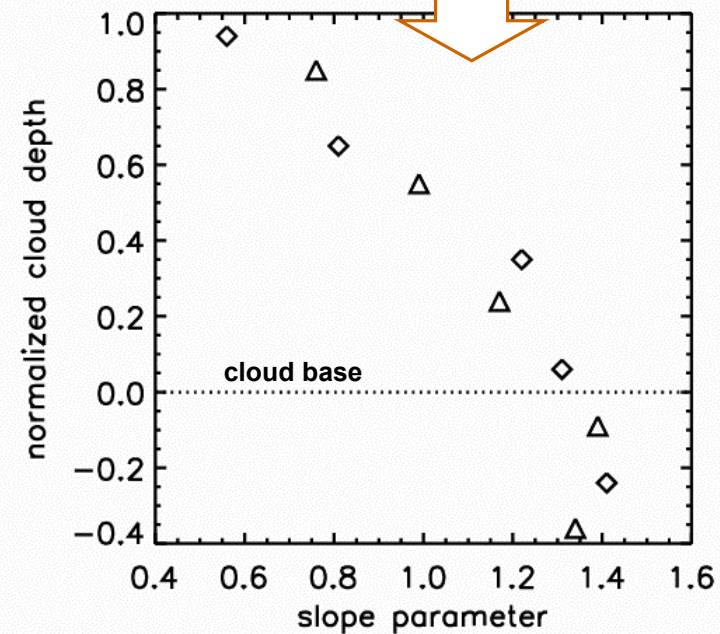
on the other hand ...



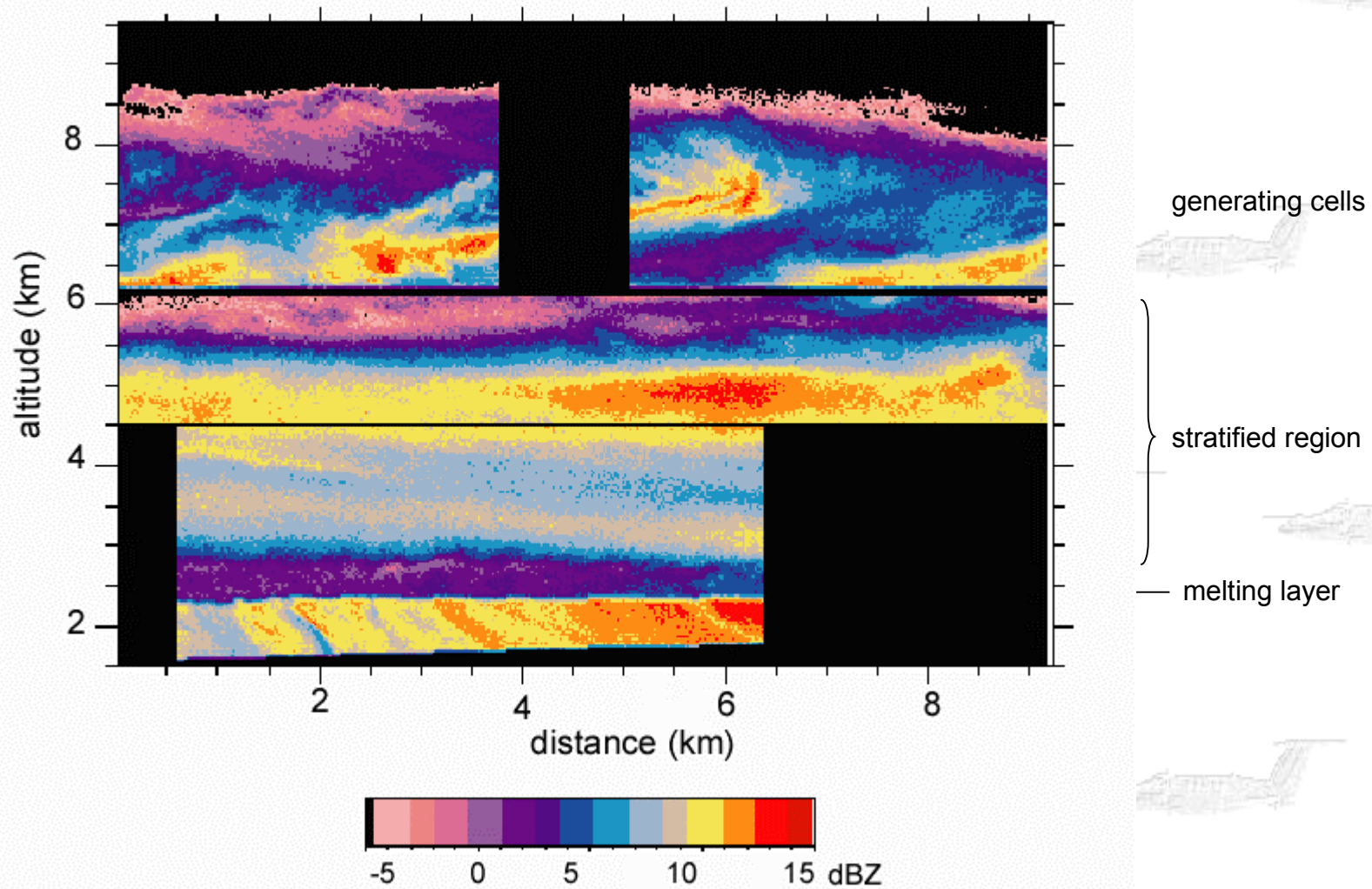
reflectivity and precipitation rate  
show strong correlations, justifying  
the use of power-law equations, like

$$Z = 4.8 R^{0.97}$$

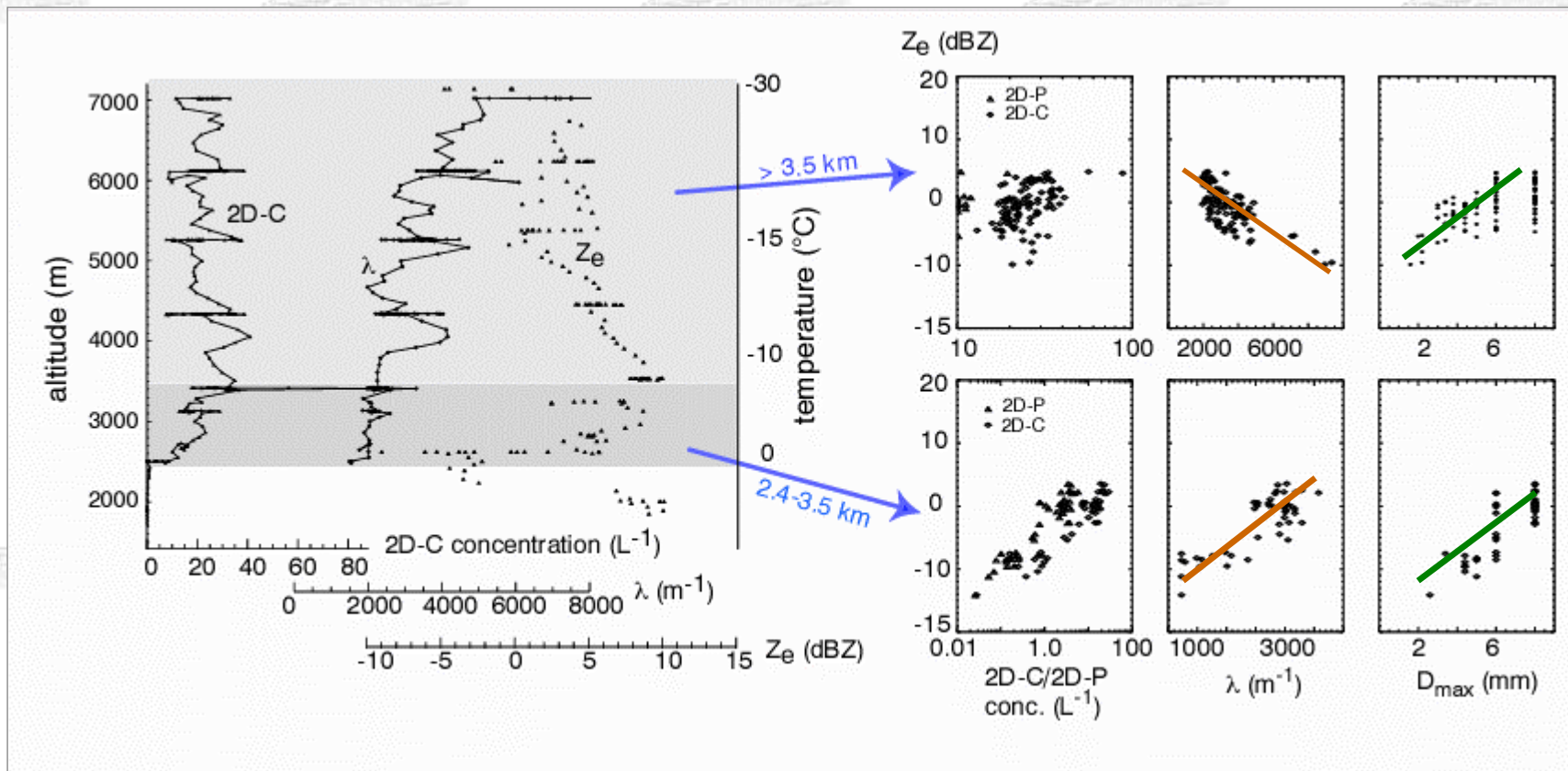
but slope parameter is altitude dependent.



950915 and 950916



October 31, 1992 **Ns**, Wyoming-Nebraska border

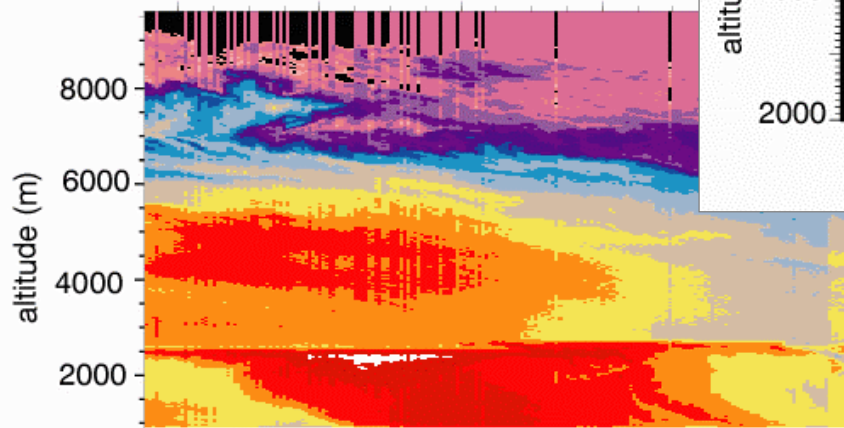
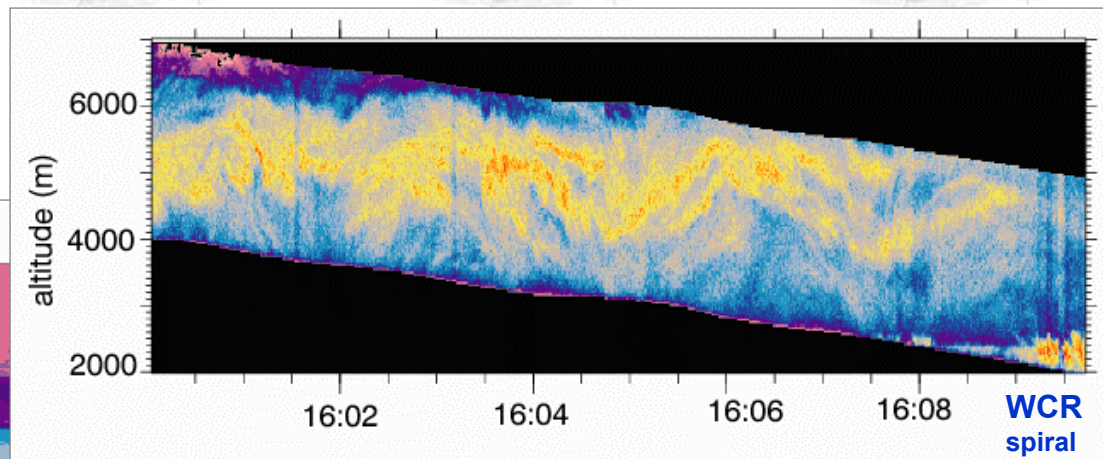


**Z vs.  $\lambda$  relationship is reversed through cloud region above the melting layer**

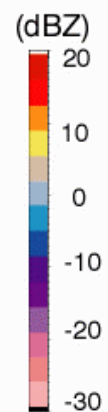
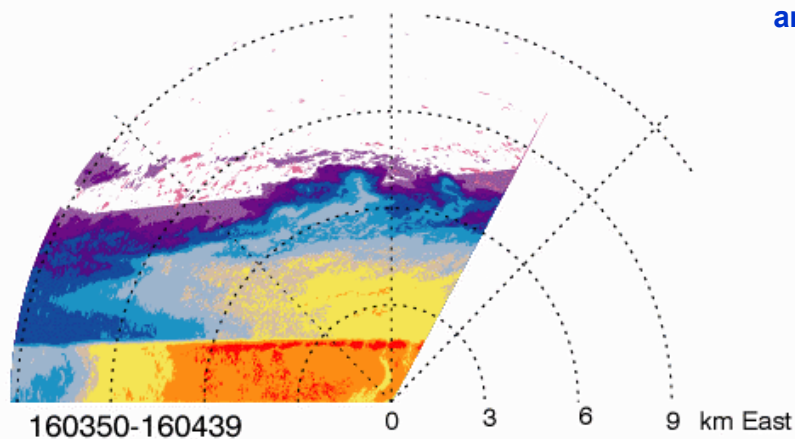
**$Z_e$  vs.  $D_{max}$  relationship appears to be more robust**

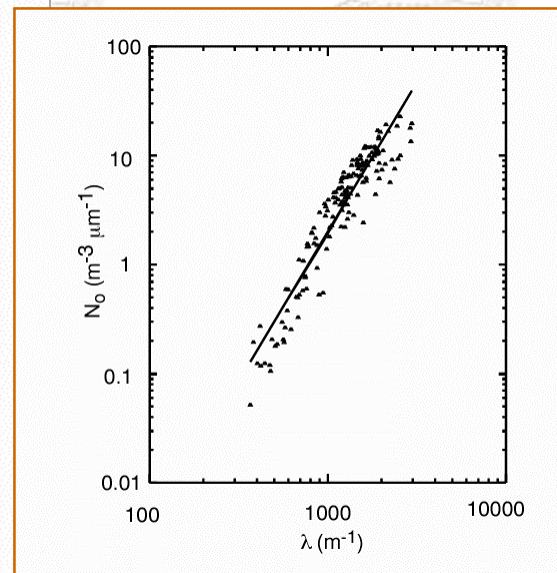
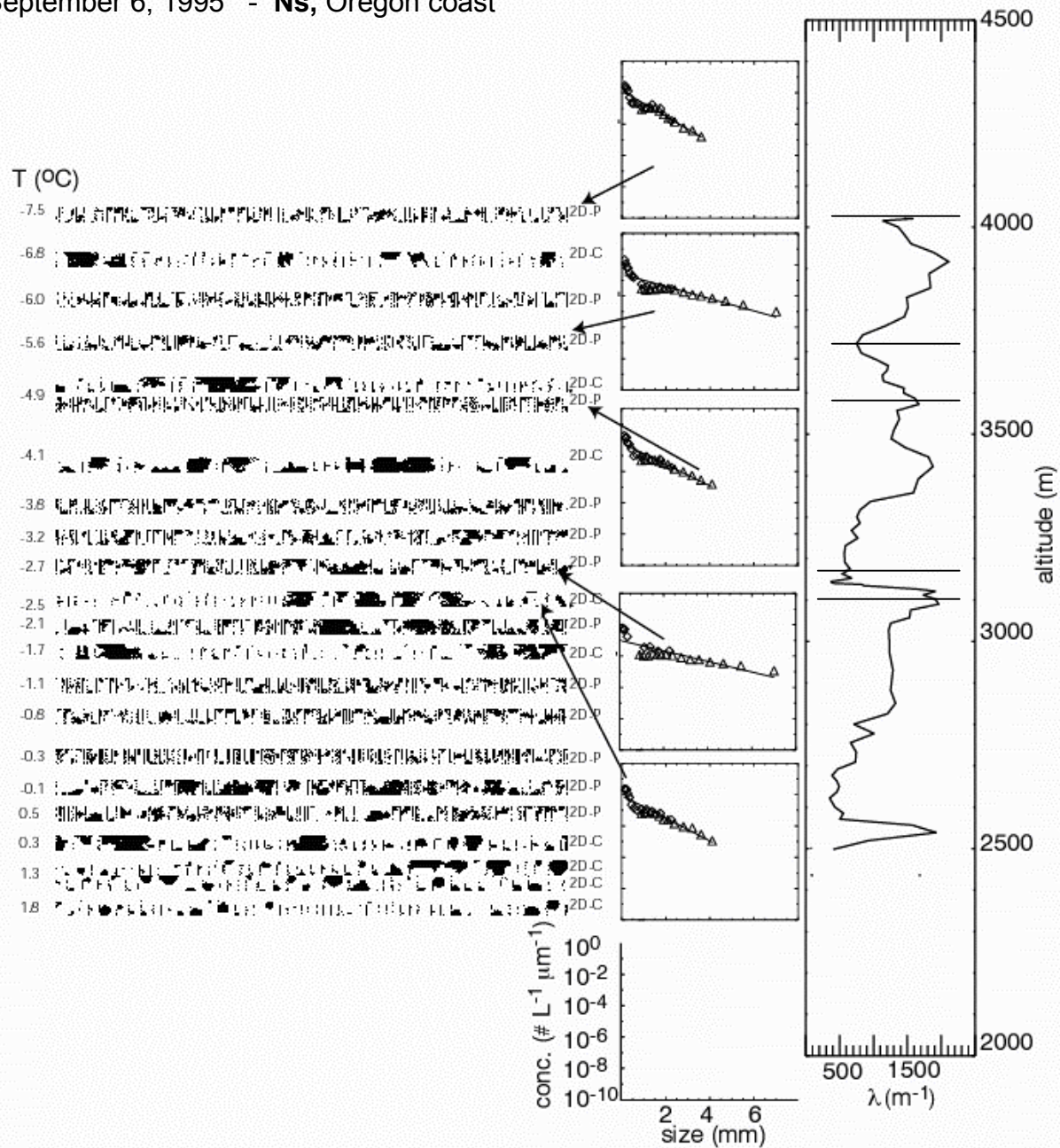


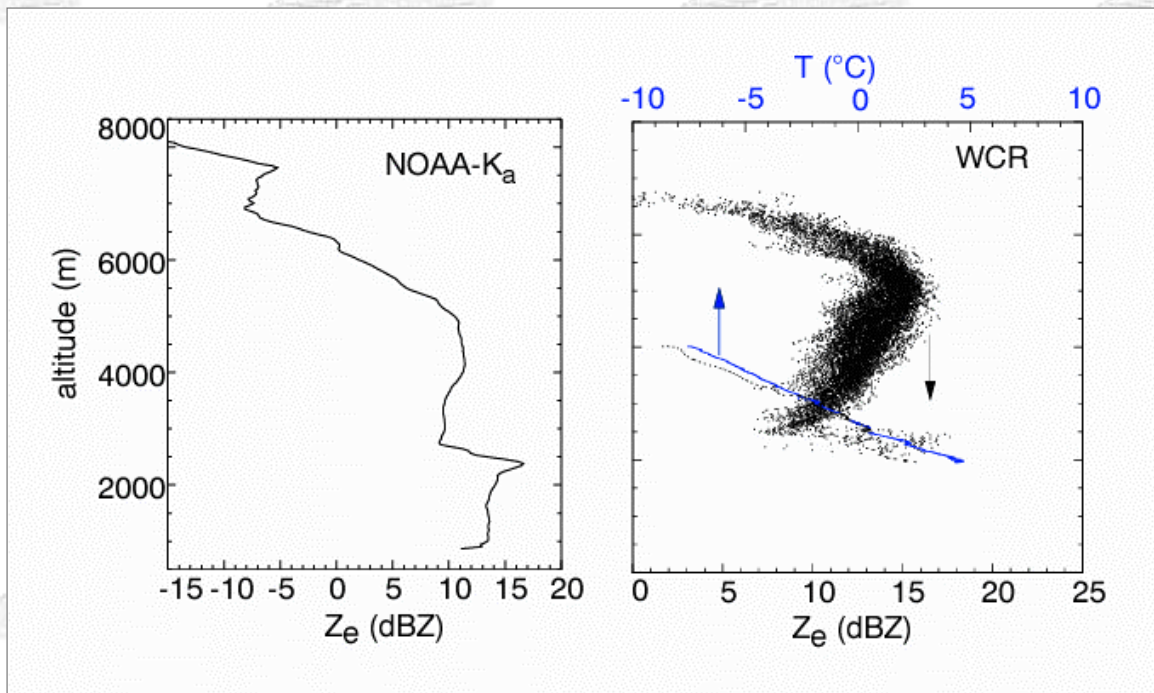
September 6, 1995 - **Ns**, Oregon coast



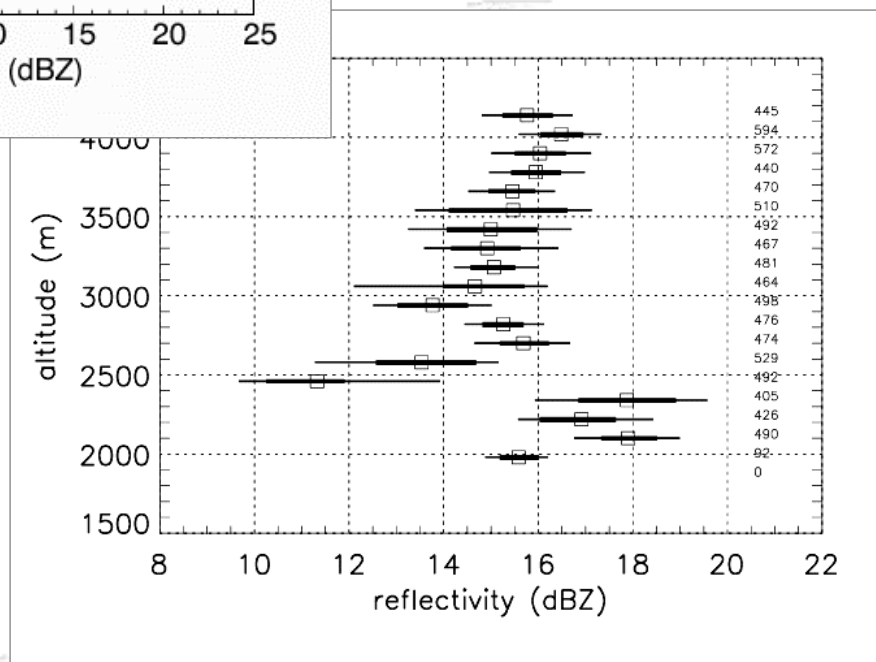
NOAA - K<sub>a</sub>  
vert. point  
and RHI



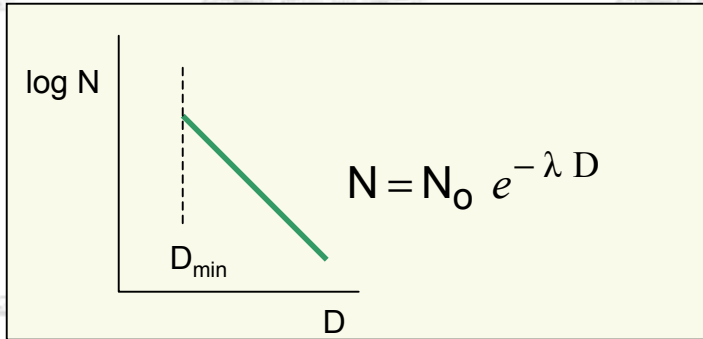
September 6, 1995 - **Ns**, Oregon coast



September 6, 1995 - **Ns**, Oregon coast



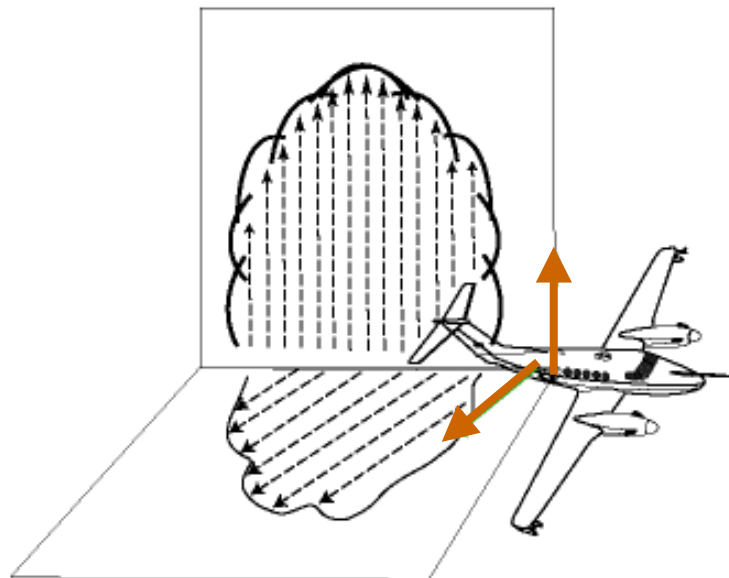




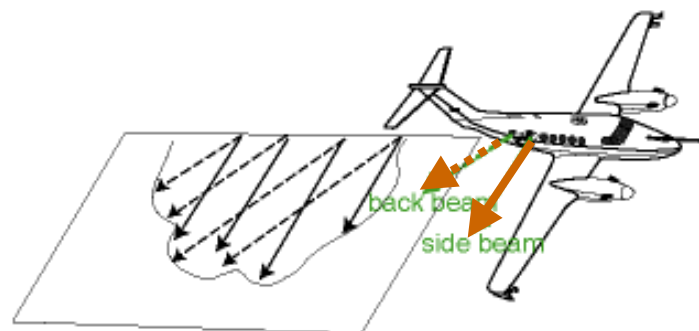
$$N_o = a \left( \frac{\lambda}{1000} \right)^b$$

Case/author	$D_{\min}$ (: m)	a	b	r	$\lambda$ ( $10^3$ ) $m^{-1}$
September 06, 1995 (Ns over Oregon)	100	3.3	2.74	0.93	3.7 - 30
October 31, 1992 (Ns over Wyoming)	100	1.88	2.32	0.91	1.5 - 9.3
October 31, 1992 (Ns over Wyoming)	900	0.68	2.41	0.91	1.2 - 3.7
Altostratus cloud (Field 1999)	800	1.54	2.89	0.7	1.9 - 5.8
Winter cyclonic storm (Lo and Passarelli 1982)	300	1.07	1.85		1.0 - 8.0
Rainband (Gordon and Marwitz 1986)	500	0.13	2.26	0.9	1.1 - 10.7

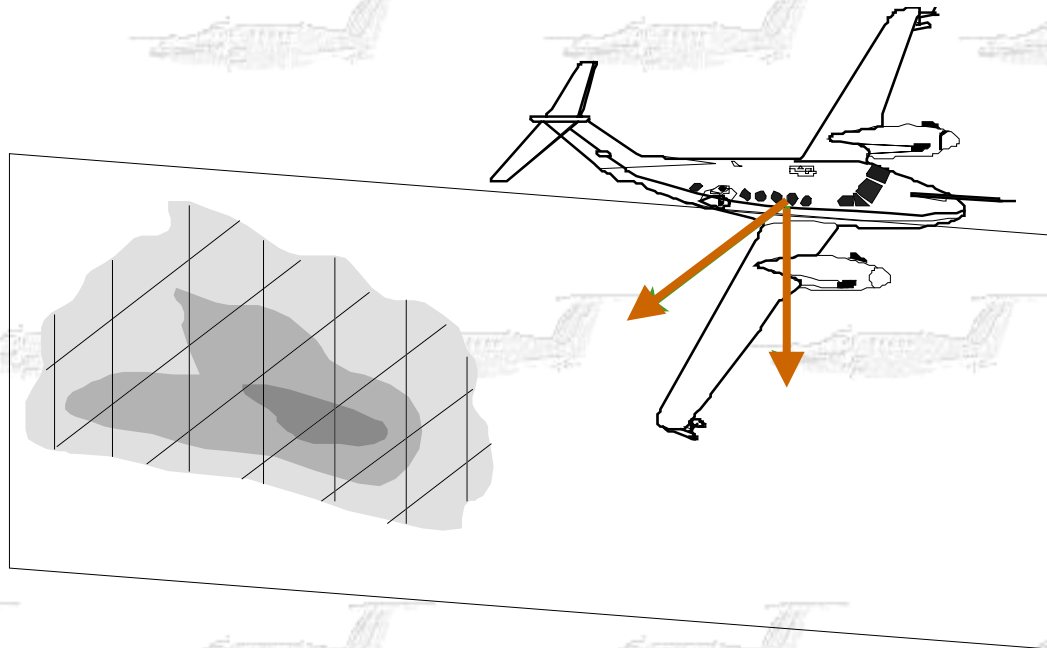
dual-direction configuration

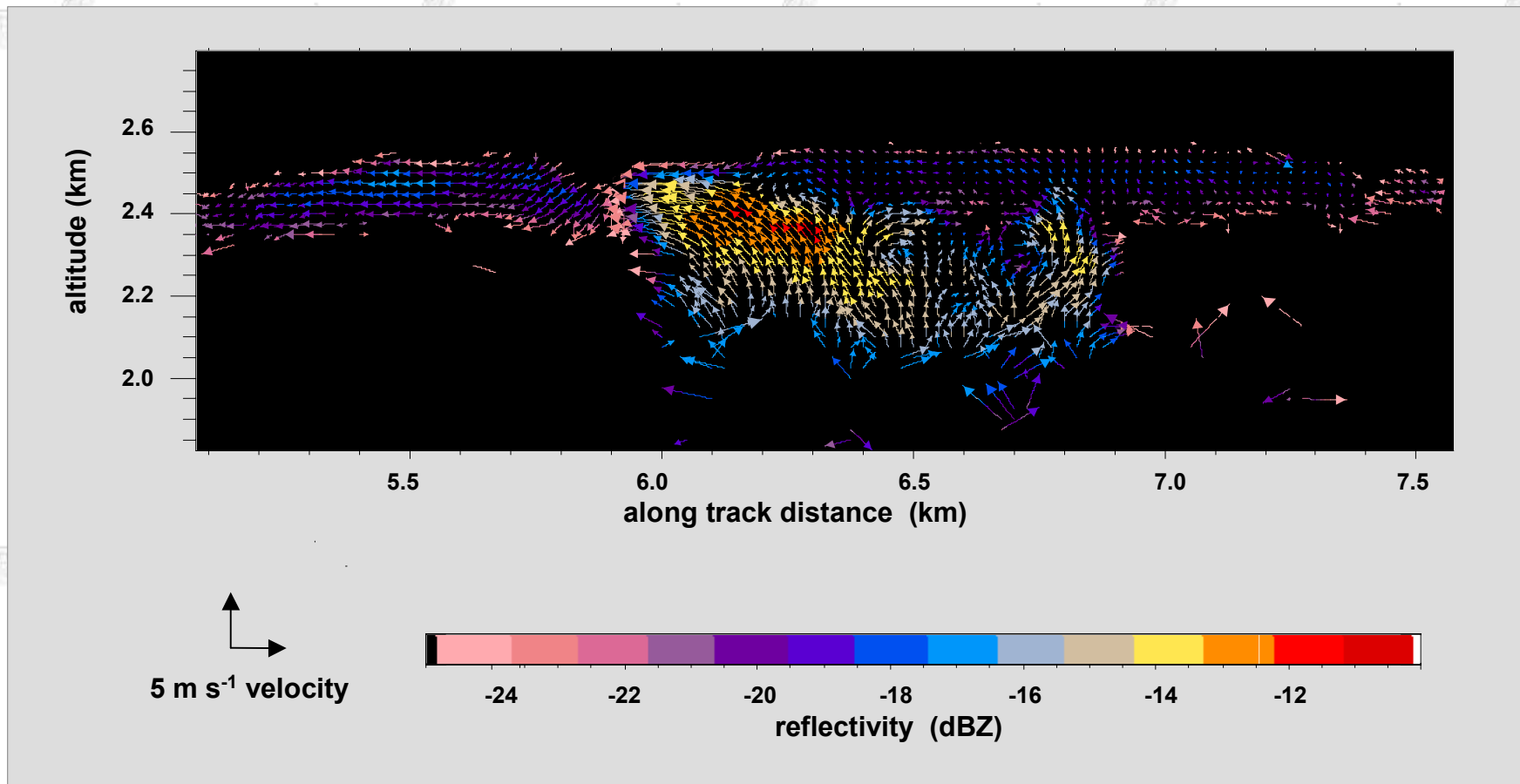


horizontal dual-beam configuration



## vertical plane dual\_Doppler





## vertical plane dual-Doppler analysis

TRAC98 980625 13:50:50 - 13:51:23